

1 **Accepted manuscript: Quaternary Science Reviews**

2 <https://doi.org/10.1016/j.quascirev.2019.105868>

3

4 **Daily climate data reveal stronger climate-growth relationships for**
5 **an extended European tree-ring network**

6 Jernej Jevšenak^{a,*}

7 *^aSlovenian Forestry Institute, Department of Forest Yield and Silviculture, Večna pot 2, 1000*
8 *Ljubljana, Slovenia*

9 *Corresponding Author: jernej.jevsenak@gozdis.si

10

11 **Abstract**

12 An extended European tree-ring network was compiled from various sources of tree-ring
13 data from Europe, northern Africa and western Asia. A total of 1860 tree-ring chronologies
14 were used to compare correlation coefficients calculated with aggregated day-wise and
15 month-wise mean temperature, sums of precipitation and standardised precipitation-
16 evapotranspiration index (SPEI). For the daily approach, climate data were aggregated over
17 periods ranging from 21 to 365 days. Absolute correlations calculated with day-wise
18 aggregated climate data were on average higher by 0.060 (temperature data), 0.076
19 (precipitation data) and 0.075 (SPEI data). Bootstrapped correlations are computationally
20 expensive and were therefore calculated on a 69.4 % subset of the data. Bootstrapped
21 correlations indicated statistically significant differences between the daily and monthly
22 approach in approximately 1 % of examples. A comparison of time windows used for

23 calculations of correlations revealed slightly later onset and earlier ending day of the year
24 for the daily approach, while the largest differences between the two approaches arise from
25 window lengths: Correlations calculated with day-wise aggregated climate data were
26 calculated using fewer days than the monthly approach. Differences in the onset and ending
27 dates of periods for the daily and monthly approaches were greater for precipitation and
28 SPEI data than for temperature data.

29

30 **Keywords:** tree rings; dendroclimatology; tree-ring network; daily climate data; climate-
31 growth relationships; *dendroTools*

32

33 **1. INTRODUCTION**

34 In dendroclimatology, various tree-ring proxies are usually compared to gridded or observed
35 station climate data with monthly resolution to analyse climate-growth relationships (Cook
36 and Kairiukstis, 1992). Monthly climate data are more easily accessible, available for most
37 land territories and have longer time spans than daily data, but at the cost of accuracy,
38 particularly when dealing with precipitation data (Hofstra et al., 2009; Yin et al., 2015). All
39 monthly data, whether they are gridded or from station observations, are derived from daily
40 climate station observations, which are the raw climate products, and then aggregated into
41 monthly datasets. In addition to the many daily observations available from the KNMI
42 Climate Explorer (<https://climexp.knmi.nl/start.cgi>), various reforecast project
43 collaborations have resulted in high quality gridded daily data, such as E-OBS gridded daily
44 datasets for Europe (Cornes et al., 2018), Berkeley Earth temperature datasets

45 (<http://berkeleyearth.org>) and various datasets provided by the National Oceanic and
46 Atmospheric Administration of the United States
47 (<https://www.esrl.noaa.gov/psd/data/gridded/tables/daily.html>).

48 Daily climate data is well integrated into various process-based models, such as the VS
49 model (Anchukaitis et al., 2006; Shishov et al., 2016) and MAIDENiso (Danis et al., 2012).
50 Some previous dendroclimatological studies have used daily climate data. Land et al. (2017)
51 reported increasing correlations between ring widths and precipitation if heavy
52 precipitation events are excluded from the precipitation data. Their study showed that the
53 annual radial growth of oak trees is mainly affected by daily precipitation sums of less than
54 10 mm. Schönbein et al. (2015) reconstructed summer precipitation based on subfossil oak
55 tree-ring data and daily precipitation records from southern Germany, while Pritzkow et al.
56 (2016) combined the earlywood vessel area of *Quercus robur* and daily temperature data
57 from northern Poland to reconstruct minimum winter temperatures back to 1810. Climate-
58 growth relationships using daily climate data have been calculated by various authors (e.g.
59 Castagneri et al., 2015; Liang et al., 2013; Sanders et al., 2014; Sun and Liu, 2016). One of
60 the first software programs for dendroclimatological studies based on daily climate data was
61 CLIMTREG, provided by Beck et al. (2013), while Jevšenak and Levanič (2018) recently
62 presented the dendroTools R package, which is designed for the R environment (R Core
63 Team, 2019) and provides various options for analysis of climate-growth relationships on
64 daily and monthly scales.

65 Combining tree-ring networks with gridded climate data can provide comprehensive spatio-
66 temporal information related to tree growth and climate sensitivity. Compiled large-scale
67 tree-ring networks have already been used for various purposes, e.g. to analyse climate-

68 growth associations for northern hemisphere tree-ring width records (St. George, 2014); to
69 evaluate the climate sensitivity of model-based forest productivity estimates (Babst et al.,
70 2013); to identify climatic drivers of global tree growth (Babst et al., 2019); to characterise
71 relationships between climate, reproduction and growth (Hackett-Pain et al., 2018); to
72 simulate radial tree growth with the VS-Lite model on a global scale (Breitenmoser et al.,
73 2014); to assess global tree-mortality (Cailleret et al., 2017); and to quantify the drought
74 effect on tree growth as a measure of vitality (Bhuyan et al., 2017). Zhao et al. (2019)
75 analysed representatives and biases of tree-ring records in the Global Tree-Ring Databank
76 (ITRDB), identified priority sampling areas and corrected identified issues, while Babst et al.
77 (2018) discussed challenges and opportunities related to tree-ring networks. No tree-ring
78 network has so far been used to analyse climate-growth relationships for daily data and to
79 compare daily and monthly climate-growth relationships. To do so, an extended European
80 tree-ring network was established using freely available data from various sources and
81 combining these data with gridded daily climate data, i.e. E-OBS daily data on a 0.1-degree
82 regular grid.

83 In this study, I compare climate-growth correlations calculated from aggregated daily and
84 monthly data of mean temperature, sums of precipitation and standardised precipitation-
85 evapotranspiration indices (SPEI). Climate data with daily resolution enable greater
86 flexibility in the analysis of climate-growth relationships and provide higher explained
87 variance in calibration models for climate reconstructions. In areas where the time period
88 related to the climate signal starts/ends near the 15th day of the month, a daily approach
89 should provide significantly greater differences between correlations calculated from day-
90 wise and month-wise aggregated climate data. An important benefit of using a daily
91 approach is the possibility to study changes in time windows over time. While the temporal

92 stability of monthly data usually enables the study of only the changes in correlation
93 coefficients over time, a daily approach enables the study of changes in temporal windows
94 over time as well. Hypothetically, this information could be used to model the divergence of
95 climate-growth relationships (Loehle, 2009) and changes in growing season patterns
96 (Linderholm, 2006). Finally, studying climate growth correlations using day-wise aggregated
97 climate data could improve our understanding of the climate signal in tree rings and enable
98 us to more accurately predict future growth under different climate scenarios. The goal of
99 this study is to highlight the advantages of using daily rather than monthly data and, at the
100 same time, expose possible caveats related to the daily approach.

101 The paper is structured as follows: in section 3.1 I give a general description of the extended
102 European tree-ring network, while correlations calculated with day-wise and month-wise
103 aggregated climate data are compared in sections 3.2 and 3.3. The time periods related to
104 the calculated correlation coefficients for the daily and monthly approach are compared in
105 section 3.4. Finally, in section 3.5 the potential applications and future extensions of the
106 daily approach are discussed. In the conclusions the main results are summarised, and
107 possible caveats of the daily approach are discussed.

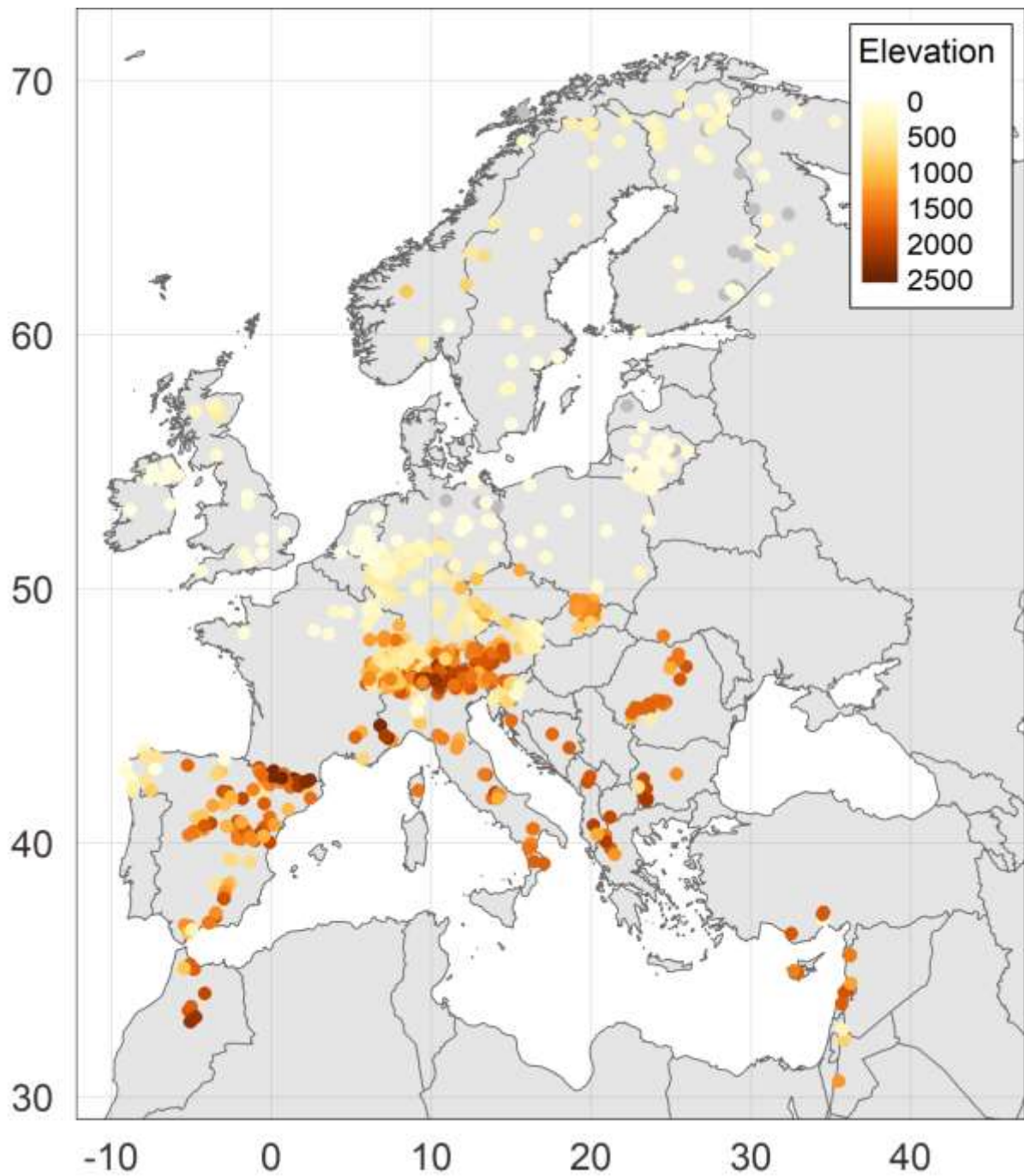
108

109 **2. MATERIALS AND METHODS**

110 **2.1 Tree-ring network**

111 For the purposes of this study, I compiled a continental-scale tree-ring network consisting of
112 freely available data from various online sources. A cleaned and corrected version of the
113 International Tree Ring Data Bank (Grissino-Mayer and Fritts, 1997), i.e. rITRDB, which was

114 presented by Zhao et al. (2019) and is available via the web repositories of the National
115 Climatic Data Center (<https://www.ncdc.noaa.gov/paleo/study/25570>), was used as the
116 primary source. This dataset consists of 8326 individual files in Tucson format, containing
117 information on various tree-ring parameters. Firstly, rITRDB was updated with 16 additional
118 rwl files from Europe, which were recently added to the International Tree Ring Data Bank.
119 These data are marked “ITRDB_2019” in Supplementary Table S1, while the files available in
120 rITRDB are marked “rITRDB”. All files were filtered to keep only those that correspond to the
121 extent of the ensemble version of the E-OBS temperature and precipitation datasets (Cornes
122 et al., 2018). E-OBS datasets cover 25°W to 45°E longitude and 25°N to 71.5°N latitude on a
123 0.1-degree regular grid (see below for a more detailed description of the E-OBS datasets).
124 Three additional filters were applied to available rwl files: all had to have at least 10 trees
125 within the site, rbar greater than 0.10 and cover at least 30 years in the period 1950 – 2018,
126 which is the time span of the E-OBS dataset. After filtering these data, I added three
127 additional datasets from the information system PANGAEA (<https://www.pangaea.de/>). One
128 dataset, presented by Tejedor et al. (2017), consists of individual measurements of conifers
129 from the Iberian Peninsula, while two datasets (Sánchez-Salguero et al., 2017; Sánchez-
130 Salguero et al., 2018) are available as already developed and standardised tree-ring width
131 chronologies. These files are marked “PANGAEA” in the source column of Supplementary
132 Table S1. Next, 521 standardised ring-width chronologies provided by Babst et al. (2013) as
133 supplementary material were added. Finally, 43 isotope chronologies, available via the
134 repository of freely available data from the BACI H2020 project ([https://www.bgc-
135 jena.mpg.de/geodb/projects/Data.php](https://www.bgc-jena.mpg.de/geodb/projects/Data.php)), were added. The final network of tree-ring data
136 consisted of 1860 chronologies (Figure 1).



137

138 **Figure 1: Locations of analysed chronologies with respective elevation. Missing elevations**
139 **are marked in grey.**

140

141 All files that were available as raw data, i.e. data from rITRDB, ITRDB and some of the
142 PANGEA files, were detrended using a spline with a 50% frequency cutoff response at 32
143 years. For detrending, I used the *detrend()* function from the dplr R package (Bunn, 2008).
144 Detrended measurements were averaged to create a single composite series describing site
145 chronology. Chronologies from PANGEA, Babst et al. (2013) and isotope chronologies from
146 the BACI repository were available as already developed and standardised series and were
147 used as such. Babst et al. (2013) used the same detrending method as in our study, while
148 chronologies from PANGEA were standardised using negative exponential or linear
149 functions (Sánchez-Salguero et al., 2018) and negative exponential or linear functions and
150 30-year-long splines (Sánchez-Salguero et al., 2017). For a description of the final database,
151 see section 3.1; the complete meta-data is available in Supplementary Table S1.

152

153 **2.2 E-OBS daily climate data and SPEI calculation**

154 To create climate variables, the daily mean, minimum and maximum air temperature and
155 sums of precipitation data were downloaded as netCDF files from
156 http://surfobs.climate.copernicus.eu/dataaccess/access_eobs.php. The E-OBS version 19.0e
157 on a 0.1-degree regular grid was used, which was released in March 2019 and covers the
158 time span from January 1st, 1950 to December 31st, 2018. Using the *knnLookup()* function
159 from the SearchTrees R package (Becker, 2012), the closest grid point was located in the E-
160 OBS dataset for each individual site, and climate data were extracted. To study climate
161 growth relationships on daily and monthly scales, mean temperature and sums of
162 precipitation data were used. In addition, downloaded minimum and maximum air

163 temperature data were used to calculate daily and monthly SPEI series (Beguería and
164 Vicente-Serrano, 2017; Vicente-Serrano et al., 2010).

165 SPEI combines precipitation data and potential evapotranspiration (*PET*) data. To calculate
166 *PET*, the Hargreaves-Samani method (Hargreaves and Samani, 1985) was used (Eq. 1), where
167 T_{mean} is mean daily air temperature, T_{max} is maximum daily air temperature, T_{min} is minimum
168 daily air temperature and R_a is net radiation at the surface ($\text{MJm}^{-2} / \text{day}$).

$$169 \quad PET = 0.0023 (T_{mean} + 17.8) \sqrt{T_{max} - T_{min}} R_a \text{ (Eq. 1)}$$

170 The R_a for each location and day was estimated from the solar constant and solar
171 declination. A more detailed procedure is described by Wang et al. (2015). Next, the climatic
172 water deficit (*D*) was calculated for each day (*i*) as the difference between the daily sum of
173 precipitation (*P*) and daily *PET* (Eq. 2). The calculated D_i values were then aggregated at
174 different daily and monthly time scales into a log-logistic probability distribution to obtain
175 the SPEI index series, following the same procedure used in the SPEI R package (Vicente-
176 Serrano et al., 2010).

$$177 \quad D_i = P_i - PET_i \text{ (Eq. 2)}$$

178

179 **2.3 Analysis of daily and monthly climate-growth relationships**

180 Climate-growth relationships for day-wise and month-wise aggregated data were analysed
181 for all chronologies in the final tree-ring network for mean temperatures, sum of
182 precipitation and SPEI. For daily data, the *daily_response()* function from the dendroTools R
183 package (Jevšenak and Levanič, 2018) was used. The *daily_response()* function works by
184 sliding a moving window through daily climate data and calculating correlation coefficients

185 between aggregated daily climate data and the selected tree-ring chronology. With the
186 *daily_response()* function, all correlation coefficients for time windows between 21 and 365
187 days were calculated. The analysis started with day of year (DOY) 1 and finished with DOY
188 365. To exclude ecologically impossible effects, e.g. the effect of individual days and very
189 short intervals on annual tree-ring parameters, the shortest time window considered was 21
190 days. Therefore, to calculate the first day-wise aggregated correlation coefficient, climate
191 variables were aggregated from DOY 1 to DOY 21, to calculate the second correlation
192 coefficient, climate variables were aggregated from DOY 2 to DOY 22, etc. The last
193 correlation coefficient was calculated using a window size of 365 days, where climate
194 variables were aggregated from DOY 1 to DOY 365. Using this approach, for each chronology
195 and climate variable, the number of calculated correlation coefficients sums to 59 685. The
196 pros and cons of this approach and possible calculations of spurious correlations are
197 discussed later in the conclusions.

198 Daily datasets were next aggregated into monthly datasets and used in the
199 *monthly_response()* function of the dendroTools R package. This function resembles the
200 idea of *daily_response()*: it calculates all possible correlation coefficients between the
201 selected tree-ring chronology and aggregated monthly climate data. All correlation
202 coefficients are therefore calculated for individual months, starting with January, as well as
203 combinations of consecutive months, starting with two consecutive months and finishing
204 with twelve consecutive months. For the monthly approach, the number of calculated
205 correlations between each climate variable and tree-ring chronology is 78. For the daily and
206 monthly approach, the correlation coefficients were calculated using the Pearson method.

207 After the calculation of all correlation coefficients with *daily_response()* and
208 *monthly_response()* functions, the highest calculated absolute correlation coefficient was
209 targeted for the daily and monthly approach and the optimal time window was defined,
210 which can be described with onset DOY, end DOY and the difference between the two, i.e.
211 the length of identified time window in days. To enable useful comparison with the daily
212 approach, the identified optimal time window for monthly data is described in DOYs. For
213 example, if the highest calculated monthly correlation coefficient was calculated for the
214 combination of the months June-July, the onset DOY was 152 (June 1st), the end was DOY
215 212 (July 31st) and the window length was 61 days.

216

217 **2.4 Data analysis**

218 All analyses were performed using R software (R Core Team, 2019). The highest calculated
219 correlations and their respective time windows from *daily_response()* and
220 *monthly_response()* were first compared, and the differences between the two were
221 analysed for all proxies together and also separately for different types of proxies. To
222 further evaluate the calculated correlations and assess the significance of the differences
223 between the daily and monthly approach, bootstrapped correlations with 1000
224 bootstrapped replicates were calculated. Bootstrapping of correlations inside
225 *daily_response()* is computationally expensive and time consuming; therefore, it was done
226 on a subsample of 69.4 % of randomly selected chronologies.

227 The meta-data of the 1860 tree-ring chronologies with calculated correlations with day-wise
228 and month-wise aggregated climate data together with related time windows are given in
229 Supplementary Table S1. Three R scripts are available via the GitHub repository

230 (https://github.com/jernejevsenak/analysis_european_tree-ring_network): File *analysis.R* is
231 executable and reproduces the main results presented in this study by using Supplementary
232 Table S1. *dendroTools.R* describes the extraction of correlations calculated with day-wise
233 and month-wise aggregated temperature and precipitation data, while *SPEI.R* describes the
234 same procedure for SPEI data. The aggregation of water balance (D_i) into daily/monthly SPEI
235 of various scales is not possible inside the *daily_response()/monthly_response()* functions
236 due to the organizational structure of both functions. Therefore, both functions were
237 modified and available in *SPEI.R*.

238

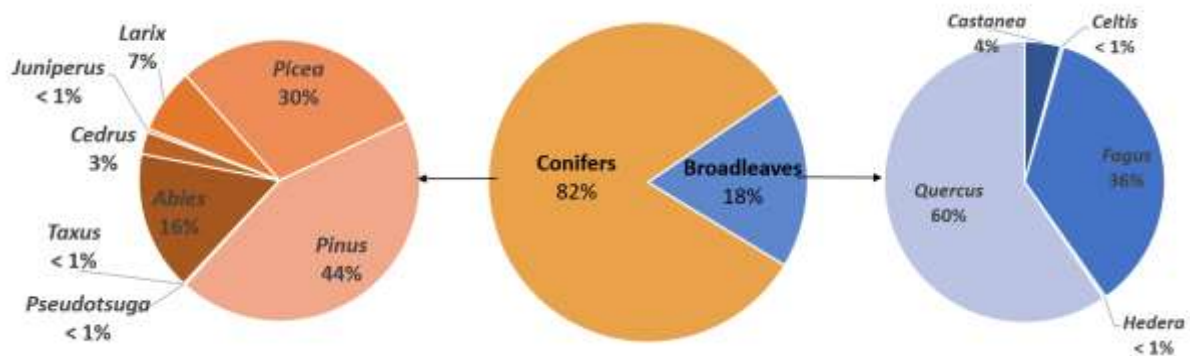
239 **3 RESULTS AND DISCUSSION**

240 **3.1 Overview of the extended European Tree-ring Network**

241 The compiled extended European tree-ring network consisted of 1860 chronologies from
242 Europe, northern Africa and western Asia, with elevations ranging from 0 to 2450 m a.s.l.
243 (Figure 1). The main contributor of the data used in this study was Fritz Schweingruber, who
244 provided 30.3 % of all files. There were 42 different tree species, with *Picea abies* being the
245 most common (445 chronologies), followed by *Pinus sylvestris* (340 chronologies), *Abies*
246 *alba* (225 chronologies), *Fagus sylvatica* (120 chronologies), *Larix decidua* (113
247 chronologies), *Pinus nigra* (101 chronologies) and *Quercus robur* (87 chronologies). The
248 majority of measurements were tree-ring widths (67.1 %), followed by early and latewood
249 measurements (6.8 % each), maximum (5.8 %) and minimum (4.7 %) density, latewood
250 percentage (2.7 %), early and latewood density (1.9 % each), stable carbon isotope ratio
251 ($\delta^{13}\text{C}$) (1.8 %) and stable oxygen isotope ratio ($\delta^{18}\text{O}$) (less than 1 %). Conifers provided 82 %
252 of analysed chronologies, with only 18 % belonging to broadleaves (Figure 2). Of the 1860

253 chronologies, 624 were available as already developed and standardised chronologies, while
 254 1236 were detrended as described in section 2.1. The mean \bar{r} of individual chronologies
 255 was 0.35, ranging from 0.10 to 0.75. The minimum number of years included in the analysis
 256 was 31, with a mean of 46 and a maximum of 67 years.

257



258

259 **Figure 2: Share of analysed data for conifers and broadleaves per genus.**

260

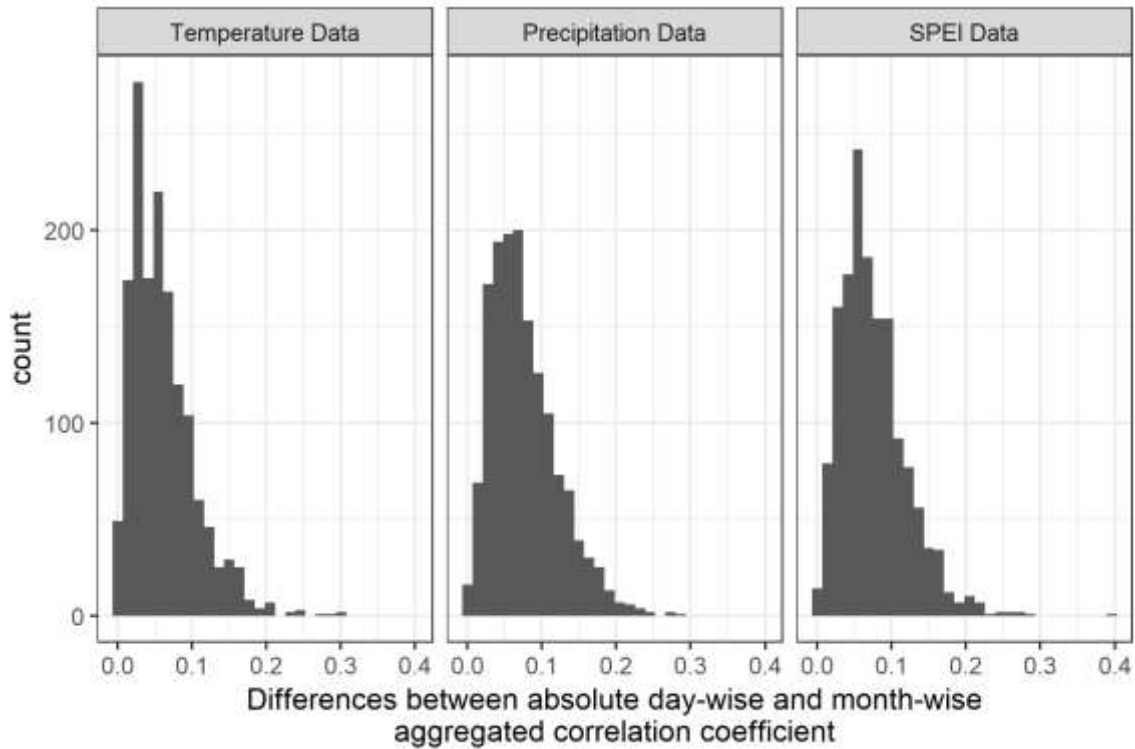
261 **3.2 Comparison of correlations calculated with day-wise and month-wise aggregated**
 262 **climate data**

263 The *daily_response()* and *monthly_response()* functions were used together with mean
 264 temperature, sum of precipitation and SPEI data on the 1860 chronologies, and the highest
 265 calculated correlation coefficients between the daily and monthly approaches were
 266 compared. While the monthly and daily approach might identify correlation coefficients
 267 from different time windows and different signs, I compared only those which had the same
 268 sign and an overlap of at least 7 days in their optimal time windows, which indicates that
 269 the correlations refer to the same climate signal. The opposite was true in about 10 % of

270 examples (10.6 % for temperature, 11.0 % for precipitation and 11.4 % for SPEI data), which
271 were not accounted for in further analysis.

272 Excluding calculations with different time windows and/or opposite positive/negative signs,
273 the mean absolute correlation coefficient with daily temperature data was 0.467, while for
274 monthly data, it was 0.407 (Table 1). On average, the correlation coefficient for day-wise
275 aggregated temperature data was 0.060 higher. A greater difference was calculated for
276 precipitation data, i.e. 0.076, for which the mean absolute daily correlation was 0.483 and
277 mean absolute monthly correlation 0.406. Similar values to precipitation data were also
278 calculated for SPEI data, where the mean absolute daily correlation was 0.485, mean
279 absolute monthly correlation was 0.410 and the difference between them was 0.075. The
280 actual benefit of using data on a daily scale could be inferred from histograms of differences
281 between the daily and monthly approach (Figure 3). In a few rare cases, the minimum
282 difference between absolute correlations calculated with day-wise and month-wise
283 aggregated data was 0, while in the most extreme case it was 0.390 (SPEI data), 0.306
284 (temperature data) and 0.286 (precipitation data). Standard deviations of differences were
285 0.043 (temperature) and 0.045 (precipitation and SPEI data).

286



287

288 **Figure 3: Differences between absolute correlations calculated with day-wise and month-**
 289 **wise aggregated temperature, precipitation and SPEI data.**

290

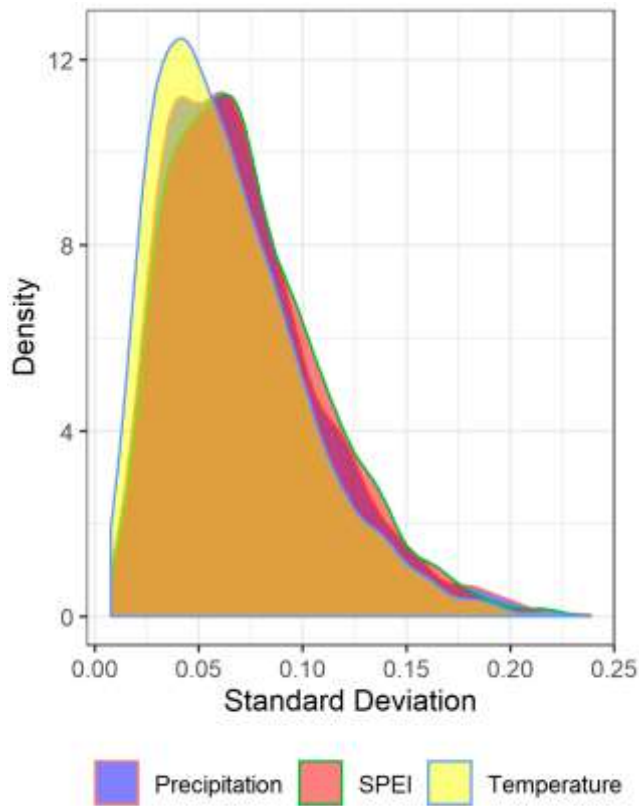
291 The observed pattern of a higher difference for precipitation and SPEI data was calculated
 292 for 8 different proxies, while for earlywood density and minimum density, the difference
 293 between correlations calculated with day-wise and month-wise aggregated data was higher
 294 for temperature data (Table 1). Calculated mean differences for proxies varied from 0.033 to
 295 0.117. Both extreme differences resulted from relatively small sample sizes. Mean
 296 differences between the daily and monthly approaches are similar to those reported by Sun
 297 and Liu (2016), who compared monthly and daily (pentated) correlations for three sites
 298 from China and obtained differences of 0.04 (precipitation), 0.06 (maximum temperature)
 299 and 0.10 (maximum temperature) in favour of daily (pentated) data. There are too few
 300 comparative studies between daily and monthly approaches to be able to compare our

301 results further. In general, standard deviations were lower for day-wise aggregated
302 correlations (Table 1), which indicates a more consistent climate signal in individual proxies.
303 The highest calculated correlation coefficients are related to maximum and latewood
304 density proxies and temperature data.

305 The benefit of using a daily rather than monthly approach is therefore greater for
306 precipitation and SPEI data. To investigate this phenomenon, the characteristics of
307 temperature, precipitation and SPEI time series were considered. Since temperature data
308 have a clear annual pattern and higher autocorrelation in comparison to precipitation and
309 SPEI data (e.g. Amirabadizadeh et al., 2015; Breinl and Di Baldassarre, 2019), the calculated
310 temperature correlation coefficient with time window X is more similar to the next one
311 calculated with the time window shifted by one day ($X + 1$ or $X - 1$). In contrast,
312 precipitation and SPEI data are less autocorrelated; therefore, change in rainfall information
313 is more rapid, and the selection of the optimal time window is of greater benefit.

314 I tested this hypothesis by comparing the variability of correlation coefficients for the three
315 climate variables. To do so, the highest calculated correlation coefficient was targeted and
316 compared with 15 previously and 15 subsequently calculated correlations. The calculations
317 had the same time window length, but were just shifted 15 days left and right on a calendar
318 scale. Then, the standard deviation of correlations was calculated and compared among
319 different climate variables. Higher standard deviations of correlation coefficients were
320 calculated for precipitation and SPEI data (Figure 4), which indicates that the change in
321 correlations is more rapid for precipitation and SPEI data in comparison to temperature
322 data. In other words, for temperature data, it matters less whether the selected time

323 window is the optimal one because shifting a few days left or right is not particularly
324 important in terms of the value of the calculated correlation coefficient.



325

326 **Figure 4: Density plot of standard deviations of calculated correlation coefficients within**
327 **the time window where the highest absolute value is located together with 15 previously**
328 **and 15 subsequently calculated correlation coefficients.**

329

330

331

332

333

334 **Table 1: Number of observations (N), mean, standard deviation, minimum and maximum**
335 **absolute daily and monthly correlation coefficients and the difference (diff) between them**
336 **for mean temperature, sum of precipitation and SPEI data. Summary statistics are given**
337 **for all proxies together as well as for different proxies separately. Calculations in which**
338 **daily and monthly correlations had different signs and/or referred to different time**
339 **periods are excluded.**

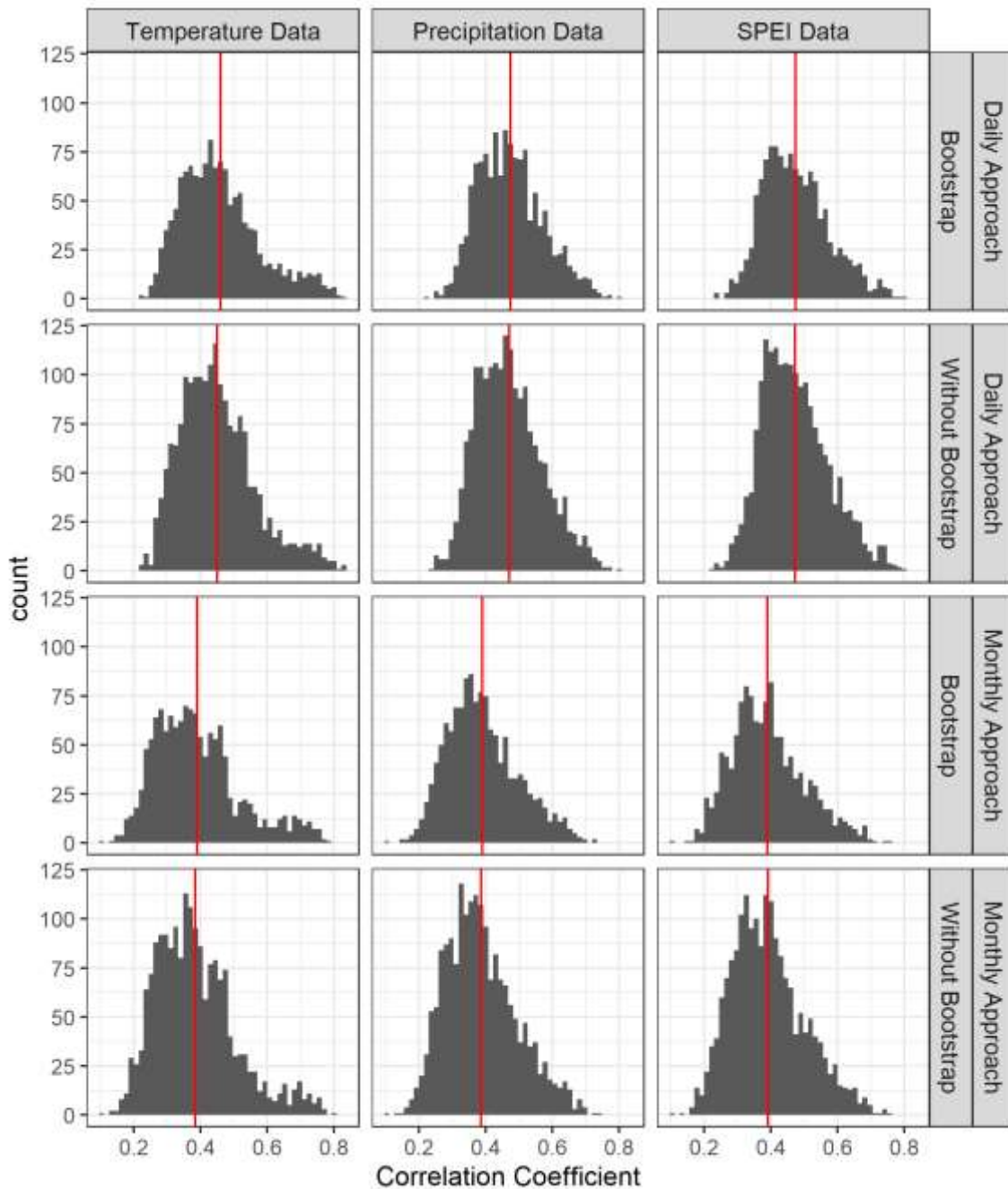
	proxy type	N	Daily Approach				Monthly Approach				diff
			mean	std	min	max	mean	std	min	max	
Temperature	All proxies	1500	0.467	0.115	0.216	0.826	0.407	0.120	0.162	0.795	0.060
	EW Density	31	0.494	0.083	0.313	0.678	0.397	0.084	0.221	0.573	0.097
	EW Width	96	0.446	0.088	0.289	0.770	0.373	0.086	0.196	0.700	0.073
	δ13C	27	0.477	0.101	0.321	0.701	0.431	0.113	0.237	0.656	0.046
	δ18O	9	0.609	0.087	0.435	0.716	0.555	0.086	0.364	0.642	0.053
	LW Density	34	0.675	0.090	0.457	0.796	0.642	0.092	0.424	0.753	0.033
	LW Percent	41	0.545	0.098	0.378	0.826	0.453	0.088	0.267	0.735	0.092
	LW Width	103	0.469	0.103	0.293	0.756	0.408	0.106	0.230	0.715	0.061
	MAX Density	100	0.664	0.104	0.317	0.826	0.625	0.111	0.247	0.795	0.039
	MIN Density	60	0.473	0.078	0.293	0.660	0.387	0.084	0.202	0.577	0.085
	Ring Width	999	0.436	0.093	0.216	0.804	0.377	0.095	0.162	0.772	0.059
Precipitation	All proxies	1500	0.483	0.099	0.250	0.802	0.406	0.106	0.159	0.735	0.076
	EW Density	22	0.474	0.063	0.387	0.636	0.385	0.083	0.272	0.594	0.089
	EW Width	107	0.455	0.085	0.269	0.661	0.374	0.090	0.163	0.597	0.081
	δ13C	33	0.518	0.110	0.338	0.704	0.462	0.127	0.260	0.668	0.056
	δ18O	3	0.417	0.049	0.361	0.454	0.345	0.049	0.291	0.388	0.072
	LW Density	31	0.524	0.085	0.342	0.710	0.448	0.104	0.233	0.652	0.076
	LW Percent	36	0.550	0.091	0.368	0.726	0.433	0.094	0.278	0.634	0.117
	LW Width	105	0.500	0.093	0.308	0.768	0.418	0.099	0.216	0.688	0.083
	MAX Density	98	0.548	0.101	0.315	0.746	0.479	0.108	0.265	0.689	0.069
	MIN Density	70	0.537	0.104	0.343	0.802	0.461	0.105	0.264	0.721	0.076
	Ring Width	995	0.469	0.095	0.250	0.764	0.394	0.103	0.159	0.735	0.075
SPEI	All proxies	1505	0.485	0.101	0.237	0.799	0.410	0.108	0.134	0.754	0.075
	EW Density	24	0.470	0.061	0.398	0.638	0.381	0.084	0.225	0.581	0.089
	EW Width	98	0.462	0.087	0.266	0.708	0.378	0.094	0.203	0.641	0.084
	δ13C	32	0.523	0.111	0.294	0.731	0.463	0.124	0.215	0.637	0.060
	δ18O	7	0.490	0.062	0.395	0.591	0.384	0.050	0.321	0.453	0.106
	LW Density	32	0.547	0.089	0.362	0.725	0.476	0.108	0.253	0.671	0.071
	LW Percent	35	0.536	0.094	0.355	0.722	0.423	0.100	0.240	0.640	0.113
	LW Width	105	0.495	0.092	0.347	0.779	0.412	0.098	0.228	0.688	0.082
	MAX Density	96	0.573	0.105	0.340	0.761	0.509	0.110	0.272	0.701	0.063
	MIN Density	70	0.538	0.111	0.326	0.799	0.454	0.116	0.264	0.748	0.084
	Ring Width	1006	0.470	0.096	0.237	0.783	0.397	0.103	0.134	0.754	0.073

340

341

342 3.3 Comparison of bootstrapped correlation coefficients

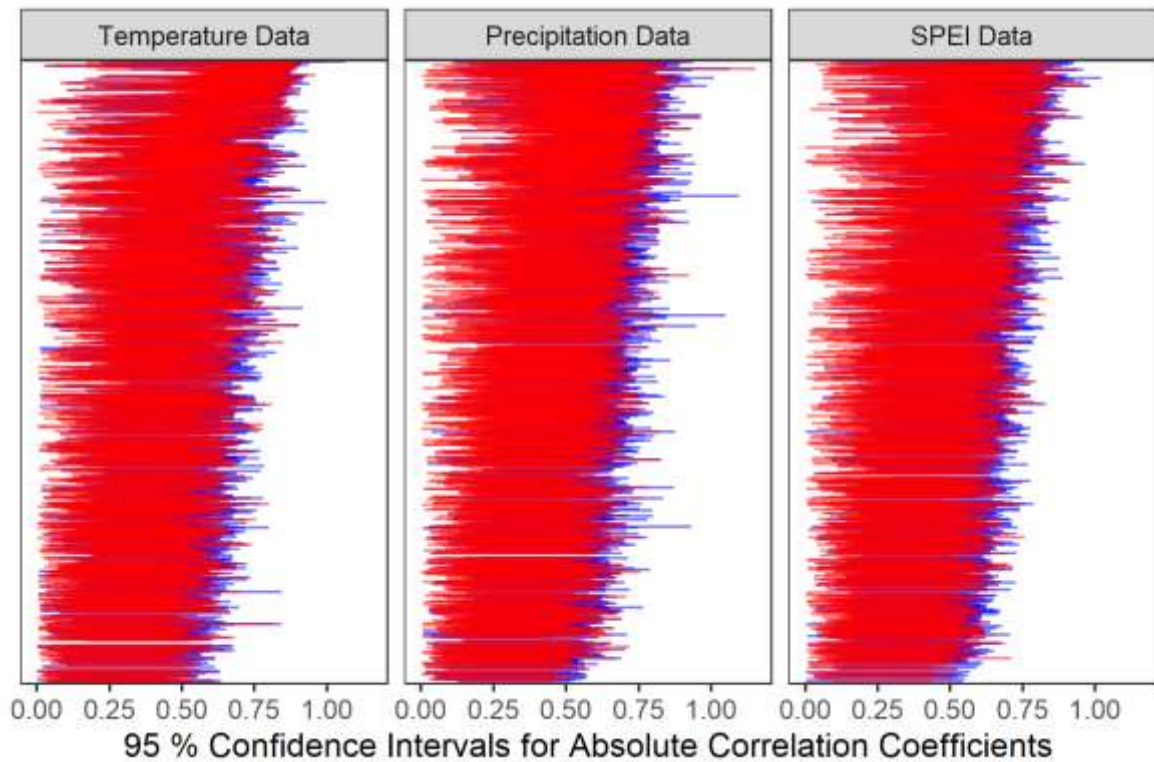
343 Due to the computationally extensive procedure, bootstrapped correlation coefficients were
344 calculated on a subset of data representing 69.4 % of all chronologies. Histograms of
345 correlation coefficients calculated with and without bootstrapping show very similar
346 patterns (Figure 5). Confidence intervals for the monthly and daily approach show a
347 considerable amount of overlap (Figure 6). To make an inference about the statistically
348 significant differences in means between the daily and monthly approach, the rule
349 presented by Cumming and Finch (2005) was used, where statistically significant differences
350 between two independent groups ($p < 0.05$) can be inferred if the share of overlap for 95 %
351 confidence intervals is no more than about half the average margin of error, that is, when
352 the proportion overlap is about 0.50 or less. The following was the case in approximately 1
353 % of examples: 0.62 % (the overlap of the daily by the monthly confidence interval) and 1.60
354 % (the overlap of the monthly by the daily confidence interval) (Figure 7). For more than 95
355 % of the calculations, the confidence intervals overlap by at least 60 %, which implies that
356 just a few examples showed statistically significant differences in means between
357 bootstrapped correlation coefficients resulting from day-wise and month-wise aggregated
358 climate data.



359

360 **Figure 5: Histograms of calculated absolute correlation coefficients for the daily and**
 361 **monthly approaches, plotted separately for the three different climate variables:**
 362 **Temperature, precipitation and SPEI data and two different strategies – with and without**
 363 **bootstrap. Red vertical lines represent the mean value of the absolute correlation**
 364 **coefficient for each group.**

365



366

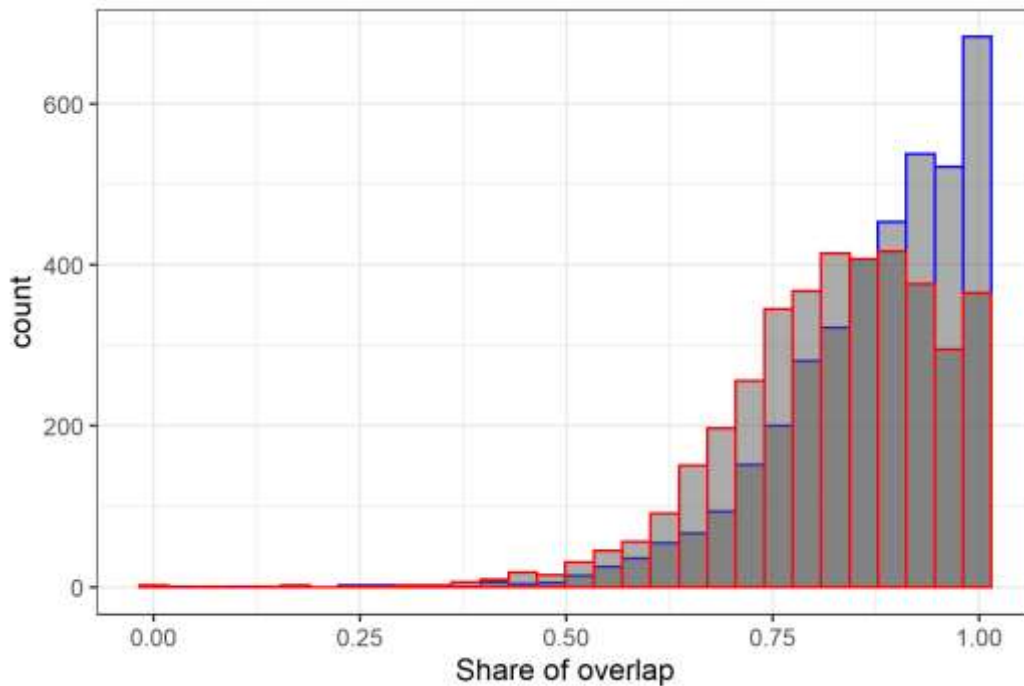
367 **Figure 6: Confidence intervals for bootstrapped correlations calculated with day-wise**
 368 **(blue) and month-wise (red) aggregated data. Only confidence intervals for correlations**
 369 **with equal signs and an overlap of at least 7 days are plotted.**

370

371

372

373



374

375 **Figure 7: Overlap of 95 % confidence intervals for bootstrapped correlations calculated**
 376 **with day-wise and month-wise climate data. The blue colour depicts the share of overlap**
 377 **of the daily by the monthly confidence interval, while the red colour depicts the share of**
 378 **the overlap of the monthly by the daily confidence interval.**

379

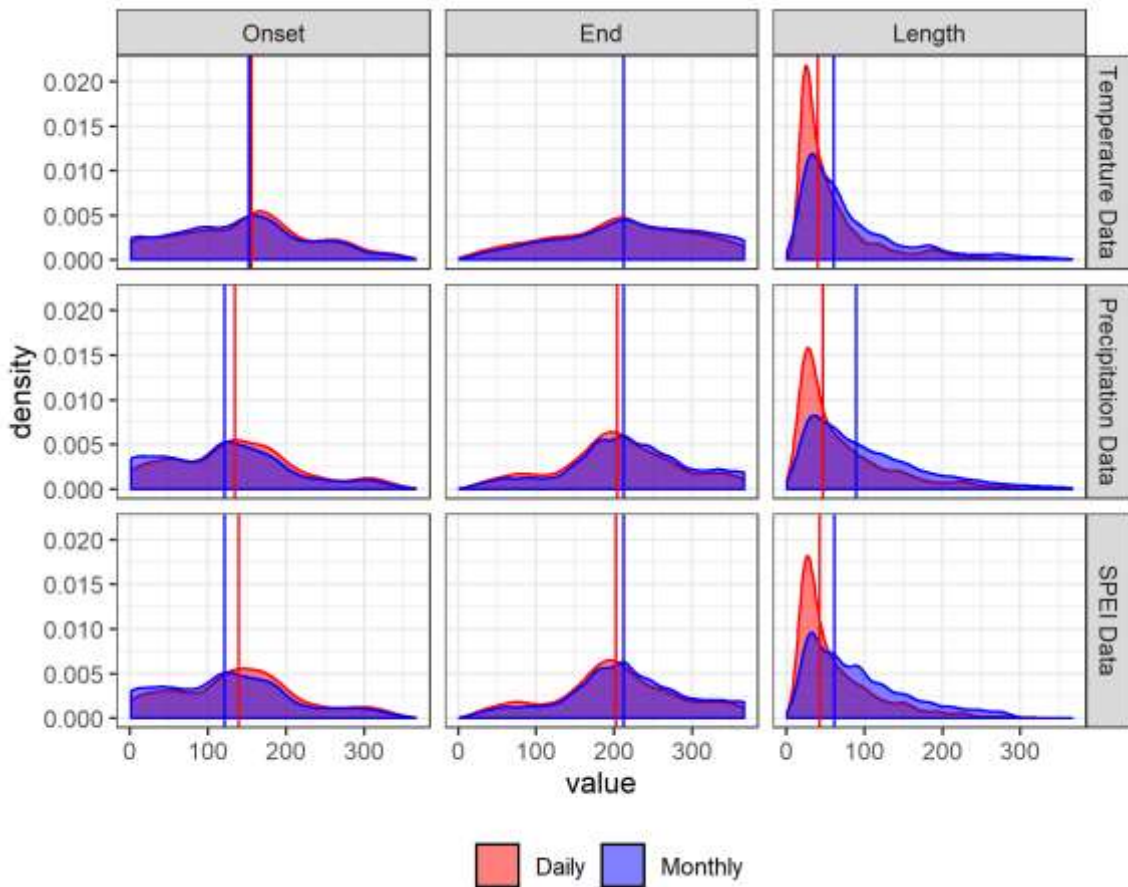
380 **3.4 Comparison of identified time windows from the daily and monthly approach**

381 Each correlation coefficient analysed in section 3.2 was calculated using a specific time
 382 window, which can be described by the onset day of year (DOY), the end DOY and the
 383 difference between them, i.e. the length of identified time windows in days. To some
 384 extent, the identified time windows are related to the growing seasons and could therefore
 385 be used to characterize growing patterns related to a specific proxy. As described in the
 386 methods section, to make a meaningful comparison between daily and monthly time
 387 windows, monthly time windows are described here in DOYs. In general, the daily approach

388 identifies later onset and earlier ending DOY (Figure 8, Table 2), while the biggest difference
389 between the two approaches arises from the window lengths: daily time windows are
390 shorter. While the median length for the daily approach was 34 (temperature), 37
391 (precipitation) and 35 (SPEI), the median length for the monthly approach was 61
392 (temperature, precipitation and SPEI data). Histograms of differences between the two
393 approaches are centred close to zero (Figure 9), indicating that the daily and monthly time
394 windows differ by a small number of days. In some rare cases, the time windows showed
395 considerable differences.

396 Similar to the comparison of correlation coefficients (section 3.2), the greatest differences in
397 time window characteristics for the daily and monthly approaches were calculated for SPEI
398 and precipitation data, while temperature data showed smaller differences between the
399 two approaches (Table 3). I assume this pattern is also related to the autocorrelation
400 present in time series of the three climate variables (see section 3.2).

401



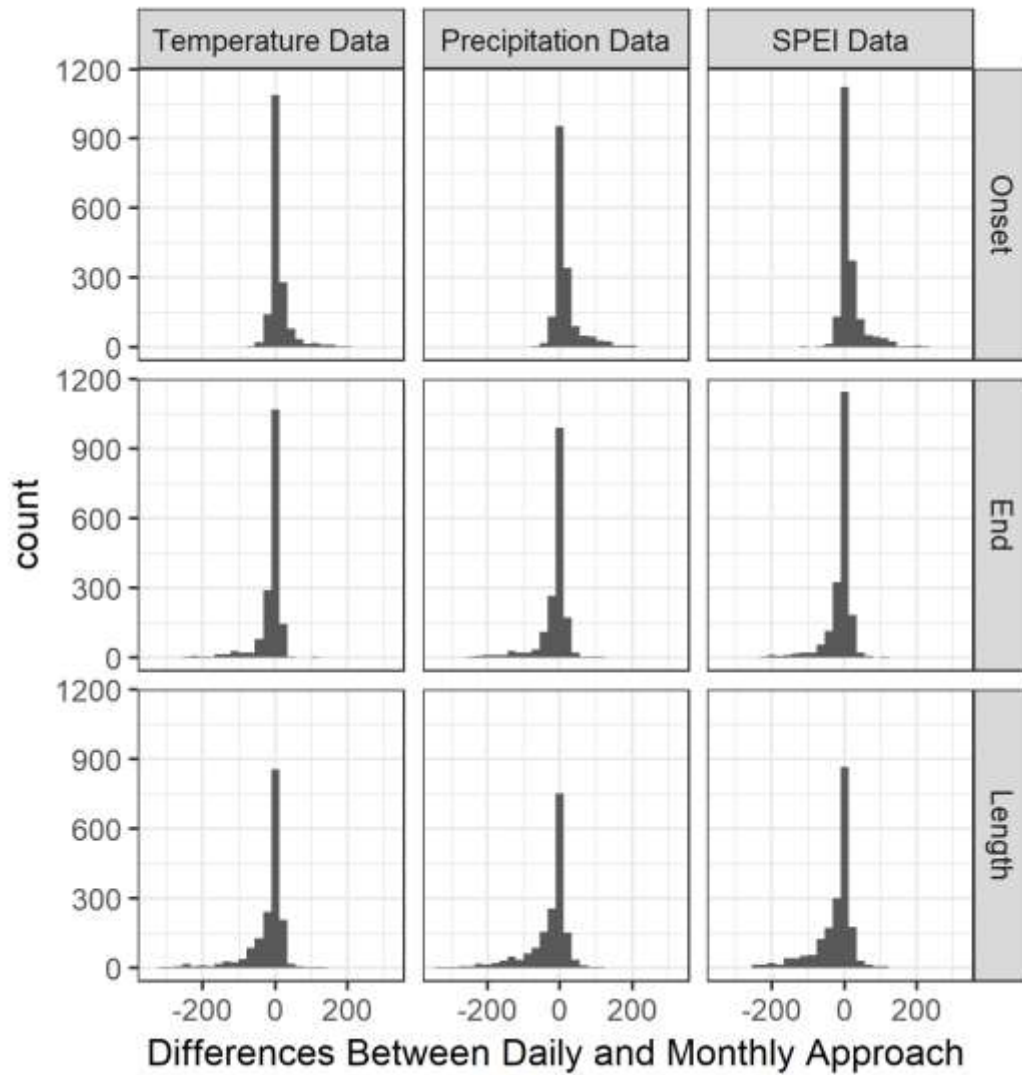
402

403 **Figure 8: Density plots for onset DOY, end DOY and time window lengths of all proxies**
 404 **plotted together for temperature, precipitation and SPEI data. Vertical lines depict**
 405 **medians.**

406 **Table 2: Median Onset DOY, End DOY and the length of optimal time windows in days, for daily and monthly approaches, separately for**
 407 **temperature, precipitation and SPEI data and different tree-ring proxies. Describing dates in brackets, given for Onset and End DOY, refer to**
 408 **a non-leap year.**

		Temperature Data			Precipitation Data			SPEI Data		
		Onset DOY	End DOY	Window Length	Onset DOY	End DOY	Window Length	Onset DOY	End DOY	Window Length
Earlywood Density	Daily	142 (May 22)	175 (Jun 24)	27	65 (Mar 06)	142 (May 22)	29	99 (Apr 09)	166 (Jun 15)	33
	Monthly	91 (Apr 01)	151 (May 31)	61	32 (Feb 01)	151 (May 31)	59	32 (Feb 01)	181 (Jun 30)	59
Earlywood Width	Daily	156 (Jun 05)	197 (Jul 16)	31	125 (May 05)	179 (Jun 28)	36	149 (May 29)	189 (Jul 08)	30
	Monthly	152 (Jun 01)	212 (Jul 31)	61	121 (May 01)	212 (Jul 31)	61	121 (May 01)	212 (Jul 31)	61
$\delta^{13}\text{C}$	Daily	158 (Jun 07)	229 (Aug 17)	32	137 (May 17)	234 (Aug 22)	58	149 (May 29)	223 (Aug 11)	56
	Monthly	152 (Jun 01)	243 (Aug 31)	62	106 (Apr 16)	258 (Sep 15)	153	91 (Apr 01)	273 (Sep 30)	184
$\delta^{18}\text{O}$	Daily	106 (Apr 16)	144 (May 24)	33	134 (May 14)	161 (Jun 10)	32	94 (Apr 04)	133 (May 13)	40
	Monthly	60 (Mar 01)	151 (May 31)	151	91 (Apr 01)	212 (Jul 31)	31	91 (Apr 01)	181 (Jun 30)	61
Latewood Density	Daily	136 (May 16)	270 (Sep 27)	89	179 (Jun 28)	242 (Aug 30)	50	177 (Jun 26)	262 (Sep 19)	64
	Monthly	121 (May 01)	273 (Sep 30)	92	182 (Jul 01)	273 (Sep 30)	92	182 (Jul 01)	273 (Sep 30)	92
Latewood Percent	Daily	160 (Jun 09)	196 (Jul 15)	27	169 (Jun 18)	217 (Aug 05)	27	191 (Jul 10)	217 (Aug 05)	27
	Monthly	152 (Jun 01)	212 (Jul 31)	31	91 (Apr 01)	212 (Jul 31)	61	121 (May 01)	212 (Jul 31)	59
Latewood Width	Daily	168 (Jun 17)	213 (Aug 01)	34	156 (Jun 05)	221 (Aug 09)	34	164 (Jun 13)	211 (Jul 30)	34
	Monthly	152 (Jun 01)	212 (Jul 31)	31	152 (Jun 01)	212 (Jul 31)	61	152 (Jun 01)	212 (Jul 31)	61
Maximum Density	Daily	185 (Jul 04)	262 (Sep 19)	71	179 (Jun 28)	250 (Sep 07)	56	180 (Jun 29)	256 (Sep 13)	57
	Monthly	182 (Jul 01)	273 (Sep 30)	92	182 (Jul 01)	273 (Sep 30)	90	182 (Jul 01)	273 (Sep 30)	92
Minimum Density	Daily	148 (May 28)	194 (Jul 13)	26	92 (Apr 02)	175 (Jun 24)	51	114 (Apr 24)	175 (Jun 24)	38
	Monthly	121 (May 01)	181 (Jun 30)	61	60 (Mar 01)	181 (Jun 30)	92	60 (Mar 01)	181 (Jun 30)	91
Ring Width	Daily	154 (Jun 03)	203 (Jul 22)	33	135 (May 15)	199 (Jul 18)	36	137 (May 17)	195 (Jul 14)	35
	Monthly	152 (Jun 01)	212 (Jul 31)	61	121 (May 01)	212 (Jul 31)	61	121 (May 01)	212 (Jul 31)	61

409



410

411 **Figure 9: Differences between the characteristics (Onset, End, Length) of identified time**
 412 **windows from the daily and monthly approach, calculated as daily minus monthly, plotted**
 413 **separately for temperature, precipitation and SPEI data.**

414

415

416

417 **Table 3: Summary statistics of differences between the characteristics (Onset, End,**
 418 **Length) of identified time windows from the daily and monthly approach, calculated as**
 419 **daily minus monthly, shown separately for temperature, precipitation and SPEI data.**

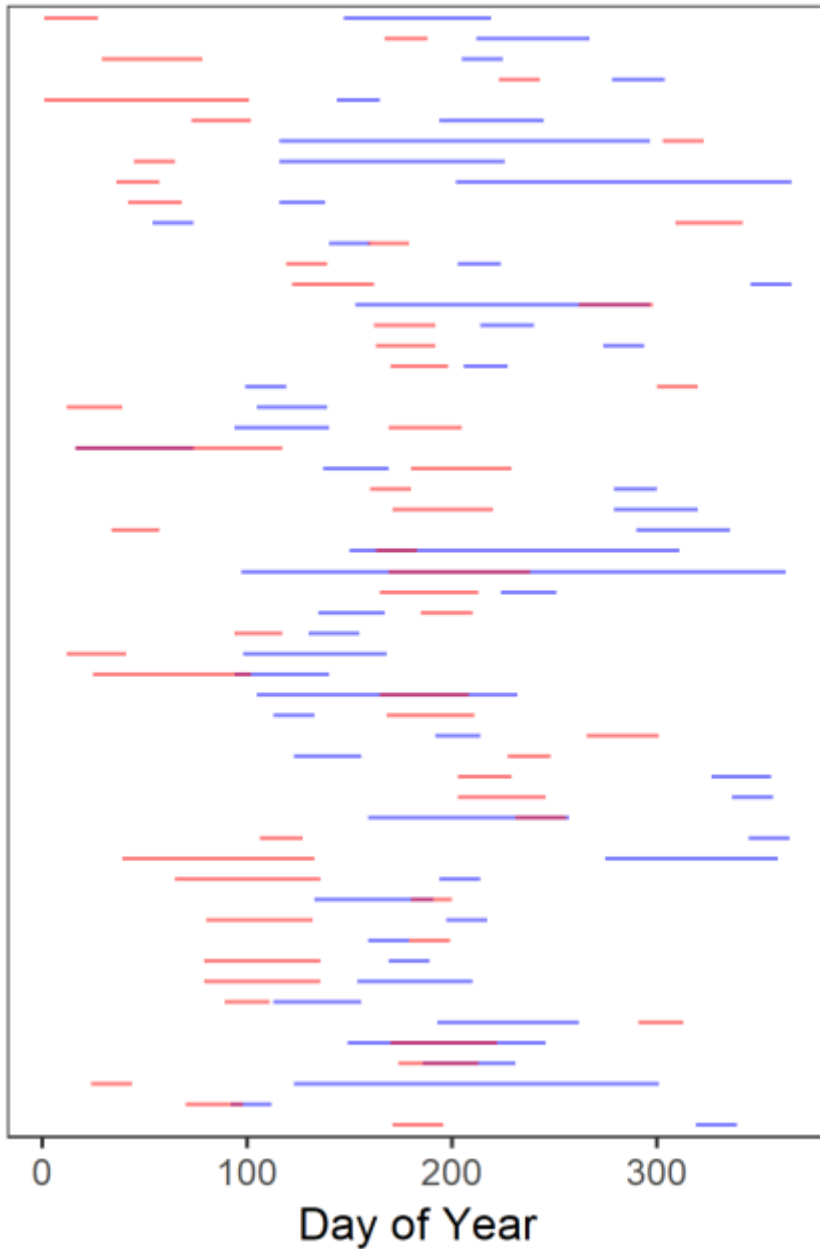
Climate variable	Window	median	std	max	min
Temperature	Onset	3	31.3	253	-169
	End	-4	38.3	107	-331
	Length	-8	55.2	183	-344
Precipitation	Onset	5	38.4	303	-165
	End	-4	42.5	184	-317
	Length	-9	63.0	247	-344
SPEI	Onset	6	35.5	254	-165
	End	-4	36.0	108	-253
	Length	-9	53.6	187	-254

420

421 **3.5 Potential applications for climate-growth relationship investigations**

422 Further dendroclimatological applications of applying climate data on a daily scale are
 423 discussed in this section. The function *daily_response()* uses flexible time windows and
 424 removes the limits defined by calendar months and therefore results in higher calculated
 425 correlation coefficients. In addition to analysis with simple and bootstrapped correlation
 426 coefficients, which was the focus in this study, the dendroTools R package also enables the
 427 calculation of partial correlations and multiproxy analysis, where instead of calculating
 428 correlations, linear or nonlinear models are fitted and afterwards (adjusted) explained
 429 variances are extracted. In this study, correlations were calculated using the Pearson
 430 method, while *daily_response()* also allows for calculations of correlations using the
 431 Spearman and Kendall methods. Finally, there are several functions available for the
 432 interpretation of calculated correlations, including plotting and summary methods, which
 433 were recently added to the dendroTools R package.

434 One of the great benefits of applying climate data on a daily scale is the ability to study the
435 changes in identified time windows over time. To illustrate this feature, I performed an
436 additional experiment in which only tree-ring width chronologies with at least 60 years of
437 data in the period 1950 – 2015 were included. Fifty-five chronologies were split into two
438 periods, i.e. early (1950 – 1980) and late (1981 – 2010 or the most recent year), and
439 analysed with the *daily_response()* function, where mean daily temperatures were used as
440 the climate variable. Identified time windows were then plotted separately for early and
441 later periods (Figure 10). Thirty-eight out of 55 (69 %) chronologies showed a shift in their
442 time windows towards earlier DOYs, which might be related to changes in growing patterns.
443 Those patterns could be investigated on a greater spatial scale, separately for different tree
444 species and elevational transects. Studying climate-growth relationships on daily scales
445 therefore opens many new possibilities. In addition to studying changes in time windows
446 over time, it would be interesting to investigate the dependences between time windows
447 and periods of the highest rate of xylem cell production. Finally, the approach from the
448 *daily_response()* could be used in various ecophysiological or climate reconstruction models,
449 where higher explained variance is expected. Following the results presented in section 3.2,
450 the daily approach on average improves the explained variance by 5.16 % (temperature
451 data), 6.65 % (precipitation data) and 6.60 % (SPEI data). Examples of climate
452 reconstructions using day-wise aggregated climate data in combination with linear and
453 nonlinear transfer functions are provided in vignettes of the *daily_response()* function on
454 CRAN [\(\[https://cran.r-\]\(https://cran.r-project.org/package=dendroTools/vignettes/Examples_daily_response.html\)](https://cran.r-project.org/package=dendroTools/vignettes/Examples_daily_response.html)
455 [project.org/package=dendroTools/vignettes/Examples_daily_response.html](https://cran.r-project.org/package=dendroTools/vignettes/Examples_daily_response.html)).



456

457 **Figure 10: Identified time windows calculated with the *daily_response()* function for early**
 458 **(blue colour) and late (red colour) periods.**

459

460 **4. CONCLUSIONS**

461 The results presented here highlight the advantages of using day-wise aggregated climate
 462 data instead of a month-wise approach. In comparison to correlations with month-wise

463 aggregated climate data, correlations with day-wise aggregated climate data were on
464 average higher by 0.060 (temperature data), 0.076 (precipitation data) and 0.075 (SPEI
465 data). The benefit of using daily data is greater for precipitation and SPEI data, while more
466 autocorrelated temperature series show smaller differences to the monthly approach. The
467 results are consistent for calculations with and without bootstrapping. Based on the share of
468 overlapped confidence intervals for bootstrapped correlations, I concluded that, except for
469 1 % of the calculations, there are no significant differences in means between day-wise and
470 month-wise aggregated correlation coefficients.

471 In this analysis, I compared only correlations which had the same sign and showed at least 7
472 days of overlap of their optimal time windows. I highlighted only the highest calculated
473 correlation coefficient resulting from the *daily_response()* and *monthly_response()*
474 functions, while potential secondary climate effects with lower correlation coefficients were
475 not considered in this study. However, all calculated correlations are saved in a matrix and
476 given as the first element of the function's output (Jevšenak and Levanič, 2018) so that
477 potential users can explore those matrices and select different time windows, if needed.
478 Furthermore, the effects of previous growing seasons are ignored solely because of
479 computational reasons. Again, *daily_response()* and *monthly_response()* can also be used to
480 analyse the effects of previous growing seasons. To do so, the argument *previous_year* must
481 be set to TRUE. The significance of the differences between the daily and monthly approach
482 were inferred from the overlap of 95 % confidence intervals. Usually, this would be done
483 with the t-test, but the calculation strategy from *daily_response()* and *monthly_response()*
484 does not allow this, since only the mean bootstrapped correlation coefficient and its lower
485 and upper confidence intervals are saved and available for comparison.

486 The *daily_response()* function comes with a much higher risk of a type I error. The number
487 of calculated correlation coefficients for each chronology and climate variable was 59 685;
488 therefore, 2984 calculations theoretically result in a type I error. For the *monthly_response()*
489 function, the number of calculated correlation coefficients is 78, where around 4
490 calculations theoretically result in type I error. In addition to much higher risk of type I
491 errors, calculations with day-wise aggregated data are time consuming, especially when
492 bootstrapping is applied. To calculate 59 685 bootstrapped correlation coefficients with
493 1000 bootstrapped samples, it takes on average slightly more than 2 hours, while for the
494 monthly approach, 78 bootstrapped correlation coefficients with 1000 bootstrapped
495 samples are calculated in around 10 – 13 seconds.

496 Despite the mostly nonsignificant differences in calculated correlation coefficients, the
497 results presented in this paper strongly encourage the tree-ring community to seriously
498 consider the use of daily climate data rather than the monthly data typically used in
499 dendroclimatological studies.

500

501 **5. ACKNOWLEDGMENTS**

502 Funding for this study was provided by the Slovene Research Agency: Program and Research
503 Group “Forest biology, ecology and technology” P4-0107 and basic research project J4-8216
504 “Mortality of lowland oak forests – consequence of lowering underground water or climate
505 change?”. I am grateful to all researchers who have uploaded their tree-ring data to various
506 online repositories. I acknowledge the E-OBS dataset from the EU-FP6 project UERRA
507 (<http://www.uerra.eu>) and the Copernicus Climate Change Service, and the data providers
508 in the ECA&D project (<https://www.ecad.eu>). The shapefile for Figure 1 was downloaded

509 from <http://www.natureearthdata.com/downloads/10m-cultural-vectors/>. Finally, many
510 thanks go to the two anonymous reviewers who suggested many improvements and ideas
511 presented in the final version of this paper.

512

513 6. REFERENCES

- 514 Amirabadizadeh, M., Huang, Y.F., Lee, T.S., 2015. Recent Trends in Temperature and Precipitation in
515 the Langat River Basin, Malaysia. *Adv. Meteorol.* 2015, 16.
- 516 Anchukaitis, K.J., Evans, M.N., Kaplan, A., Vaganov, E.A., Hughes, M.K., Grissino-Mayer, H.D., Cane,
517 M.A., 2006. Forward modeling of regional scale tree-ring patterns in the southeastern United
518 States and the recent influence of summer drought. *Geophys. Res. Lett.* 33.
- 519 Babst, F., Bodesheim, P., Charney, N., Friend, A.D., Girardin, M.P., Klesse, S., Moore, D.J.P., Seftigen,
520 K., Björklund, J., Bouriaud, O., Dawson, A., DeRose, R.J., Dietze, M.C., Eckes, A.H., Enquist, B.,
521 Frank, D.C., Mahecha, M.D., Poulter, B., Record, S., Trouet, V., Turton, R.H., Zhang, Z., Evans,
522 M.E.K., 2018. When tree rings go global: Challenges and opportunities for retro- and
523 prospective insight. *Quat. Sci. Rev.* 197, 1-20.
- 524 Babst, F., Bouriaud, O., Poulter, B., Trouet, V., Girardin, M.P., Frank, D.C., 2019. Twentieth century
525 redistribution in climatic drivers of global tree growth. *Sci. Adv.* 5, eaat4313.
- 526 Babst, F., Poulter, B., Trouet, V., Tan, K., Neuwirth, B., Wilson, R., Carrer, M., Grabner, M., Tegel, W.,
527 Levanic, T., Panayotov, M., Urbinati, C., Bouriaud, O., Ciais, P., Frank, D., 2013. Site- and
528 species-specific responses of forest growth to climate across the European continent. *Global*
529 *Ecol. Biogeogr.* 22, 706-717.
- 530 Beck, W., Sanders, T.G.M., Pofahl, U., 2013. CLIMTREG: Detecting temporal changes in climate–
531 growth reactions – A computer program using intra-annual daily and yearly moving time
532 intervals of variable width. *Dendrochronologia* 31, 232-241.
- 533 Becker, G., 2012. SearchTrees: Spatial Search Trees. R package version 0.5.2. [https://CRAN.R-](https://CRAN.R-project.org/package=SearchTrees)
534 [project.org/package=SearchTrees](https://CRAN.R-project.org/package=SearchTrees).
- 535 Beguería, S., Vicente-Serrano, S.M., 2017. SPEI: Calculation of the Standardised Precipitation–
536 Evapotranspiration Index. R package version 1.7. <https://CRAN.R-project.org/package=SPEI>.
- 537 Bhuyan, U., Zang, C., Menzel, A., 2017. Different responses of multispecies tree ring growth to
538 various drought indices across Europe. *Dendrochronologia* 44, 1-8.
- 539 Breinl, K., Di Baldassarre, G., 2019. Space-time disaggregation of precipitation and temperature
540 across different climates and spatial scales. *J. Hydrol.: Regional Studies* 21, 126-146.
- 541 Breitenmoser, P., Brönnimann, S., Frank, D., 2014. Forward modelling of tree-ring width and
542 comparison with a global network of tree-ring chronologies. *Clim. Past* 10, 437-449.
- 543 Bunn, A.G., 2008. A dendrochronology program library in R (dplR). *Dendrochronologia* 26, 115-124.
- 544 Cailleret, M., Jansen, S., Robert, E.M.R., Desoto, L., Aakala, T., Antos, J.A., Beikircher, B., Bigler, C.,
545 Bugmann, H., Caccianiga, M., Čada, V., Camarero, J.J., Cherubini, P., Cochard, H., Coyea,
546 M.R., Čufar, K., Das, A.J., Davi, H., Delzon, S., Dorman, M., Gea-Izquierdo, G., Gillner, S.,

- 547 Haavik, L.J., Hartmann, H., Hereş, A.-M., Hultine, K.R., Janda, P., Kane, J.M., Kharuk, V.I.,
548 Kitzberger, T., Klein, T., Kramer, K., Lens, F., Levanic, T., Linares Calderon, J.C., Lloret, F.,
549 Lobo-Do-Vale, R., Lombardi, F., López Rodríguez, R., Mäkinen, H., Mayr, S., Mészáros, I.,
550 Metsaranta, J.M., Minunno, F., Oberhuber, W., Papadopoulos, A., Peltoniemi, M., Petritan,
551 A.M., Rohner, B., Sangüesa-Barreda, G., Sarris, D., Smith, J.M., Stan, A.B., Sterck, F.,
552 Stojanović, D.B., Suarez, M.L., Svoboda, M., Tognetti, R., Torres-Ruiz, J.M., Trotsiuk, V.,
553 Villalba, R., Vodde, F., Westwood, A.R., Wyckoff, P.H., Zafirov, N., Martínez-Vilalta, J., 2017. A
554 synthesis of radial growth patterns preceding tree mortality. *Global Change Biol.* 23, 1675-
555 1690.
- 556 Castagneri, D., Petit, G., Carrer, M., 2015. Divergent climate response on hydraulic-related xylem
557 anatomical traits of *Picea abies* along a 900-m altitudinal gradient. *Tree Physiol.* 35, 1378-
558 1387.
- 559 Cook, E.R., Kairiukstis, L.A., 1992. *Methods of Dendrochronology: Applications in the Environmental*
560 *Sciences.* Kluwer Academic Publishers, Dordrecht.
- 561 Cornes, R.C., van der Schrier, G., van den Besselaar, E.J.M., Jones, P.D., 2018. An Ensemble Version of
562 the E-OBS Temperature and Precipitation Data Sets. *Geophys. Res. Lett.: Atmospheres* 123,
563 9391-9409.
- 564 Cumming, G., Finch, S., 2005. Inference by eye: confidence intervals and how to read pictures of
565 data. *Am. Psychol.* 60, 170-180.
- 566 Danis, P.A., Hatté, C., Misson, L., Guiot, J., 2012. MAIDENiso: a multiproxy biophysical model of tree-
567 ring width and oxygen and carbon isotopes. *Can. J. For. Res.* 42, 1697-1713.
- 568 Grissino-Mayer, H.D., Fritts, H.C., 1997. The International Tree-Ring Data Bank: an enhanced global
569 database serving the global scientific community. *The Holocene* 7, 235-238.
- 570 Hackett-Pain, A.J., Ascoli, D., Vacchiano, G., Biondi, F., Cavin, L., Conedera, M., Drobyshev, I., Liñán,
571 I.D., Friend, A.D., Grabner, M., Hartl, C., Kreyling, J., Lebourgeois, F., Levanič, T., Menzel, A.,
572 van der Maaten, E., van der Maaten-Theunissen, M., Muffler, L., Motta, R., Roibu, C.-C.,
573 Popa, I., Scharnweber, T., Weigel, R., Wilmking, M., Zang, C.S., 2018. Climatically controlled
574 reproduction drives interannual growth variability in a temperate tree species. *Ecol. Lett.* 21,
575 1833-1844.
- 576 Hargreaves, G.H., Samani, Z., 1985. Reference Crop Evapotranspiration From Temperature. *Appl.*
577 *Eng. Agric.* 1, 96-99.
- 578 Hofstra, N., Haylock, M., New, M., Jones, P.D., 2009. Testing E-OBS European high-resolution gridded
579 data set of daily precipitation and surface temperature. *Geophys. Res. Lett.: Atmospheres*
580 114.
- 581 Jevšenak, J., Levanič, T., 2018. *dendroTools*: R package for studying linear and nonlinear responses
582 between tree-rings and daily environmental data. *Dendrochronologia* 48, 32-39.
- 583 Land, A., Remmele, S., Schönbein, J., Küppers, M., Zimmermann, R., 2017. Climate-growth analysis
584 using long-term daily-resolved station records with focus on the effect of heavy precipitation
585 events. *Dendrochronologia* 45, 156-164.
- 586 Liang, W., Heinrich, I., Simard, S., Helle, G., Liñán, I.D., Heinken, T., 2013. Climate signals derived
587 from cell anatomy of Scots pine in NE Germany. *Tree Physiol.* 33, 833-844.
- 588 Linderholm, H.W., 2006. Growing season changes in the last century. *Agric. For. Meteorol.* 137, 1-14.
- 589 Loehle, C., 2009. A mathematical analysis of the divergence problem in dendroclimatology. *Clim.*
590 *Change* 94, 233-245.

591 Pritzkow, C., Wazny, T., Heußner, K.U., Słowiński, M., Bieber, A., Liñán, I.D., Helle, G., Heinrich, I.,
592 2016. Minimum winter temperature reconstruction from average earlywood vessel area of
593 European oak (*Quercus robur*) in N-Poland. *Palaeogeogr., Palaeoclimatol., Palaeoecol.* 449,
594 520-530.

595 R Core Team, 2019. R: A language and environment for statistical computing, Vienna, Austria.

596 Sánchez-Salguero, R., Camarero, J.J., Carrer, M., Gutiérrez, E., Alla, A.Q., Andreu-Hayles, L., Hevia, A.,
597 Koutavas, A., Martínez-Sancho, E., Nola, P., Papadopoulos, A., Pasho, E., Toromani, E.,
598 Carreira, J.A., Linares, J.C., 2017. Climate extremes and predicted warming threaten
599 Mediterranean Holocene firs forests refugia. *Proc. Natl. Acad. Sci. U.S.A.* 114, E10142-
600 E10150.

601 Sánchez-Salguero, R., Camarero, J.J., Rozas, V., Génova, M., Olano, J.M., Arzac, A., Gazol, A.,
602 Caminero, L., Tejedor, E., de Luis, M., Linares, J.C., 2018. Resist, recover or both? Growth
603 plasticity in response to drought is geographically structured and linked to intraspecific
604 variability in *Pinus pinaster*. *J. Biogeogr.* 45, 1126-1139.

605 Sanders, T., Pitman, R., Broadmeadow, M., 2014. Species-specific climate response of oaks (*Quercus*
606 spp.) under identical environmental conditions. *iForest* 7, 61-69.

607 Schönbein, J., Land, A., Friedrich, M., Glaser, R., Küppers, M., 2015. Seasonal reconstruction of
608 summer precipitation variability and dating of flood events for the millennium between 3250
609 and 2250 Years BC for the Main Region, southern Germany, in: Schulz, M., Paul, A. (Eds.),
610 Integrated Analysis of Interglacial Climate Dynamics (INTERDYNAMIC). Springer International
611 Publishing, Cham, pp. 127-131.

612 Shishov, V.V., Tychkov, I.I., Popkova, M.I., Ilyin, V.A., Bryukhanova, M.V., Kirryanov, A.V., 2016. VS-
613 oscilloscope: A new tool to parameterize tree radial growth based on climate conditions.
614 *Dendrochronologia* 39, 42-50.

615 St. George, S., 2014. An overview of tree-ring width records across the Northern Hemisphere. *Quat.*
616 *Sci. Rev.* 95, 132-150.

617 Sun, C., Liu, Y., 2016. Climate Response of Tree Radial Growth at Different Timescales in the Qinling
618 Mountains. *PLoS ONE* 11, e0160938.

619 Tejedor, E., Saz, M.A., Esper, J., Cuadrat, J.M., de Luis, M., 2017. Summer drought reconstruction in
620 northeastern Spain inferred from a tree ring latewood network since 1734. *Geophys. Res.*
621 *Lett.* 44, 8492-8500.

622 Vicente-Serrano, S.M., Beguería, S., López-Moreno, J.I., 2010. A Multiscalar Drought Index Sensitive
623 to Global Warming: The Standardized Precipitation Evapotranspiration Index. *J. Clim.* 23,
624 1696-1718.

625 Wang, Q., Shi, P., Lei, T., Geng, G., Liu, J., Mo, X., Li, X., Zhou, H., Wu, J., 2015. The alleviating trend of
626 drought in the Huang-Huai-Hai Plain of China based on the daily SPEI. *Int. J. Climatol.* 35,
627 3760-3769.

628 Yin, H., Donat, M.G., Alexander, L.V., Sun, Y., 2015. Multi-dataset comparison of gridded observed
629 temperature and precipitation extremes over China. *Int. J. Climatol.* 35, 2809-2827.

630 Zhao, S., Pederson, N., D'Orangeville, L., HilleRisLambers, J., Boose, E., Penone, C., Bauer, B., Jiang, Y.,
631 Manzanedo, Rubén D., 2019. The International Tree-Ring Data Bank (ITRDB) revisited: Data
632 availability and global ecological representativity. *J. Biogeogr.* 46, 355-368.

633